



Spectral Decomposition of Cloud Radiative Effect and Cloud Radiative Feedbacks

Xiuhong Chen, Xianglei Huang
The University of Michigan
Qing Yue
Caltech/JPL

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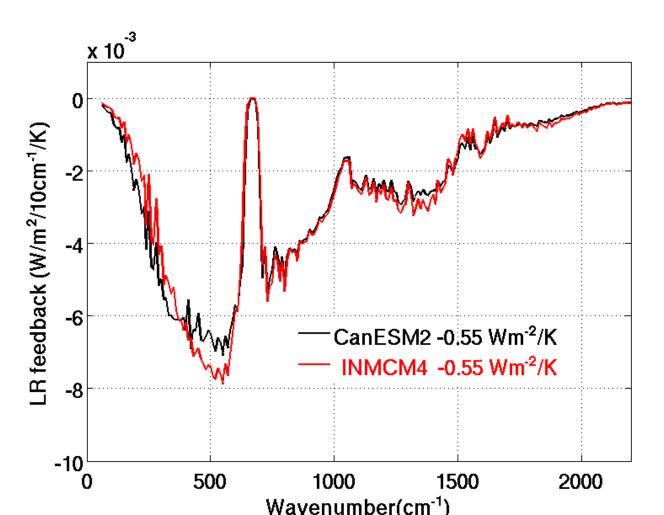
Outline

- Why go from broadband to spectral?
 - Two examples
- Traits in the spectral decomposition of CRE
- Band-by-band LW CRE: CESM vs. obs
- Band-by-band LW cloud feedbacks: 2xCO₂ CESM simulations
- Band-by-band LW short-term cloud radiative feedbacks (fluctuations): model vs. obs in 2003-2013
- Discussion and Conclusions

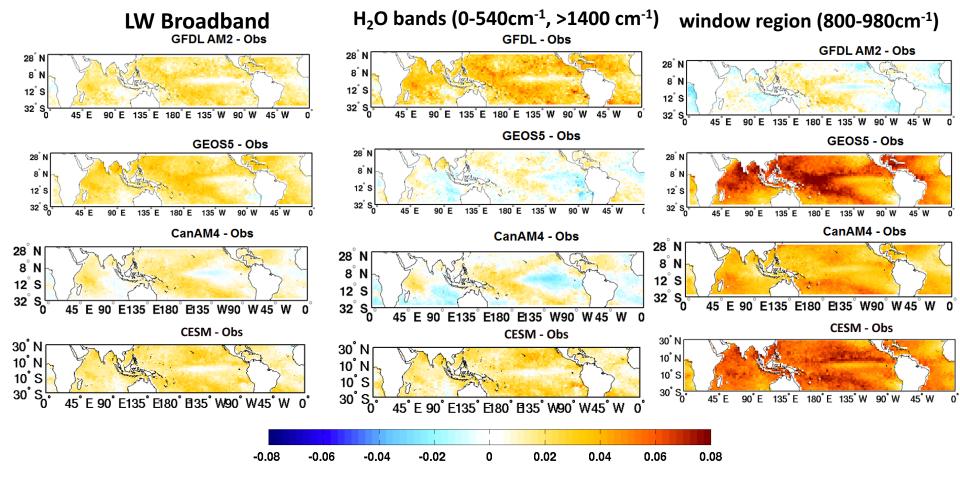


What spectral dimension can offer?

Reveal compensating differences that cannot be revealed in broadband diagnostics alone.



Spectral decomposition of broadband lapse-rate feedback (Huang et al., 2014, GRL)



clear-sky green-house efficiency

$$g_{\Delta v} = \frac{\int_{\Delta v} B_{v}(T_{s}) dv - F_{\Delta v}(TOA)}{\int_{\Delta v} B_{v}(T_{s}) dv}$$

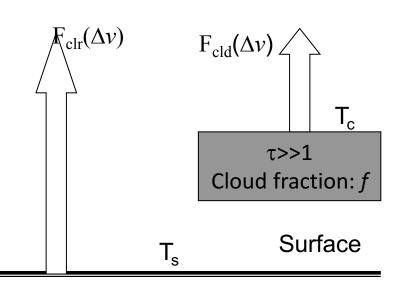
AMIP runs forced by observed SST

Obs from collocated AIRS and CERES (Huang et al., 2008; Chen et al., 2013)

(GEOS5 simulation provided by L. Oreopoulos et al; CanAM4 provided by J. Cole)



A trait of spectral (band-by-band) CRE



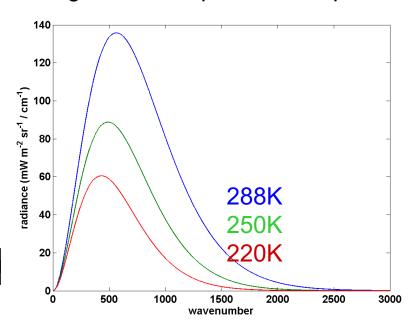
$$CRE_{LW} = \sigma T_s^4 - [f\sigma T_c^4 + (1-f)\sigma T_s^4] = f[\sigma T_s^4 - \sigma T_c^4]$$

$$CRE(\Delta v) = f[F_{clr}(\Delta v) - F_{cld}(\Delta v)]$$

Fractional contribution

$$r(\Delta v) = \frac{CRE(\Delta v)}{CRE_{LW}} = \frac{F_{clr}(\Delta v) - F_{cld}(\Delta v)}{\left[\sigma T_s^4 - \sigma T_c^4\right]}$$

- 1. Blackbody cloud
- 2. Ignore atmospheric absorption



 $r(\Delta v)$ changes with T_c

Band-to-Band ratio: sensitive to CTH but not cloud amount

LW CRE: sensitive to both CTH and cloud amount

Outcome: ratio-then-broadband approach (Huang et al., 2014, J Climate)



Derivation of spectral fluxes/CRE/feedbacks

Observations

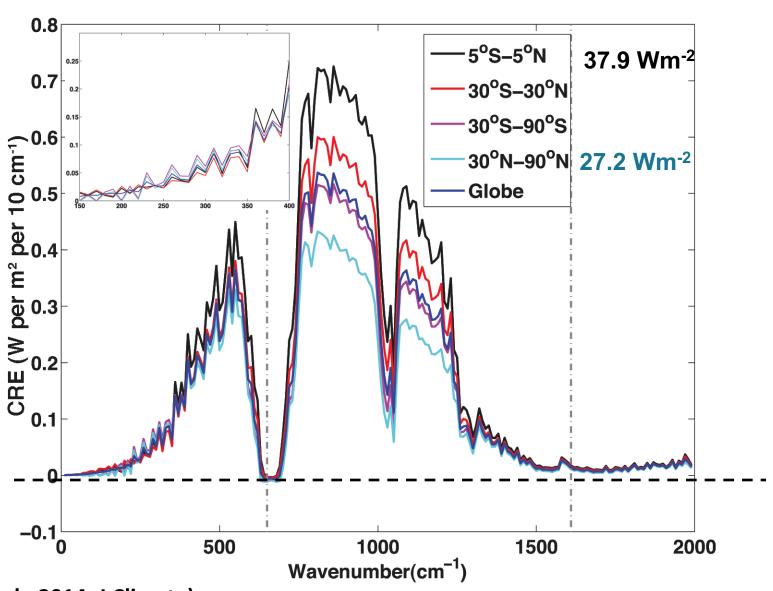
- Invert from AIRS radiances following the scene type classification of CERES (Huang et al., 2008; Chen et al, 2013; Huang et al., 2014)
- Outcome: spectral flux at 10cm⁻¹ interval over the entire LW spectrum (09/2002 to present)
- Observation-based cloud radiative kernel (Yue et al., 2016)

Models

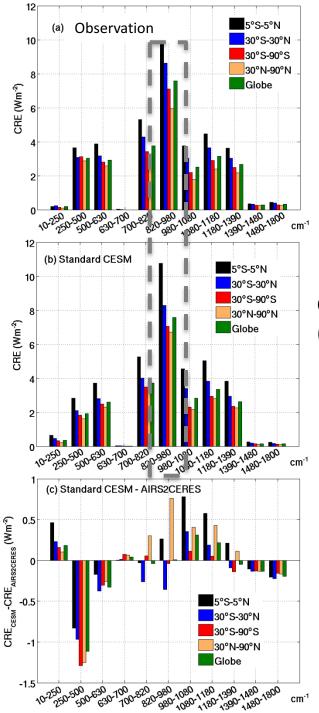
- Simple code modification to output band-by-band fluxes
- Spectral radiative kernels (Huang et al., 2014, GRL) to derive spectral details of Planck/Lapse-rate/WV feedbacks
- Adjustment methods to get spectral cloud radiative feedbacks



10-year mean spectral CRE over the different climate zones



(Huang et al., 2014, J Climate)

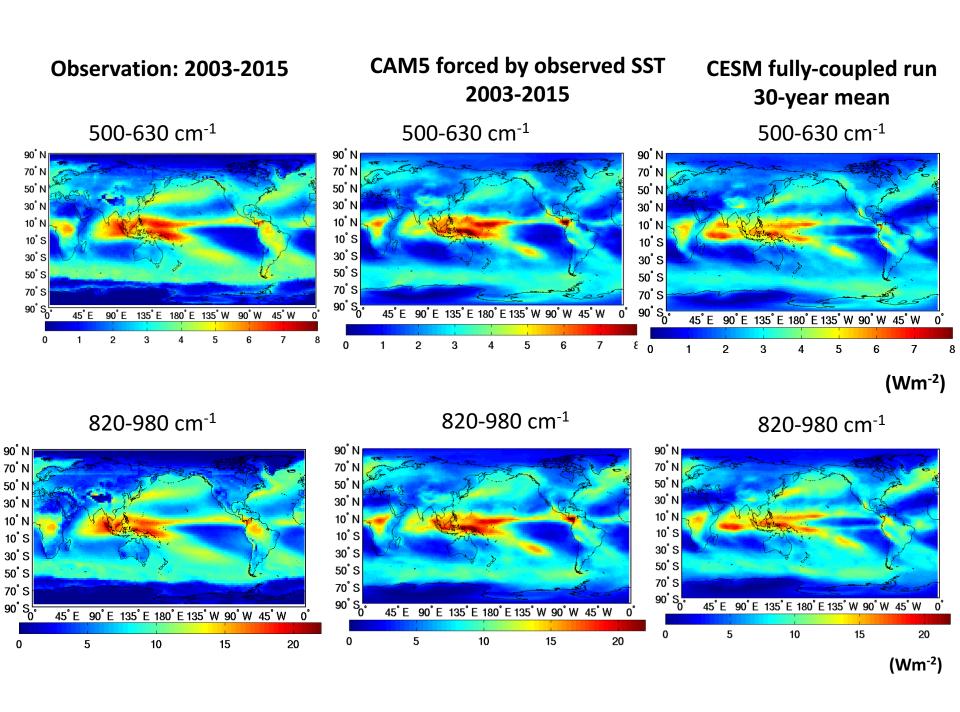


Band-by-band CRE (RRTMG_LW bandwidths)

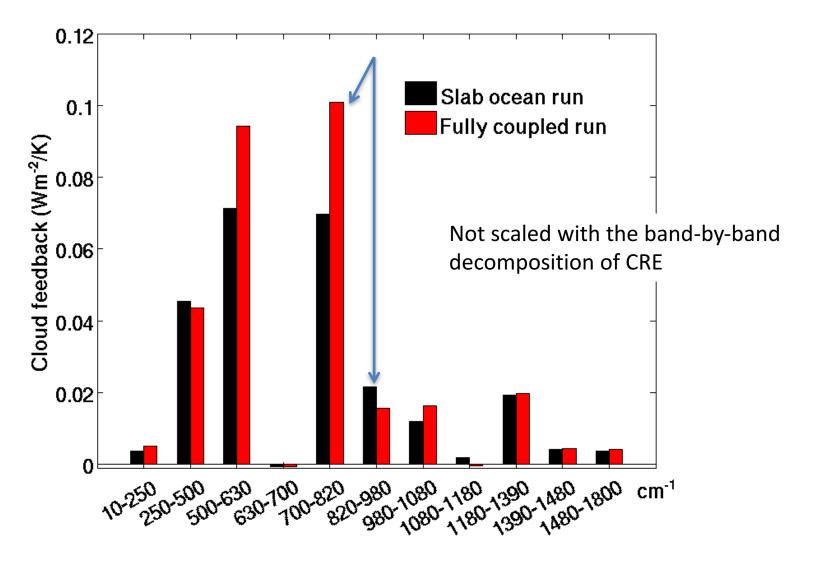
Observed averages of 2003-2015

CAM5 forced with observed SST from 2003 to 2015 (total run 2000-2015)

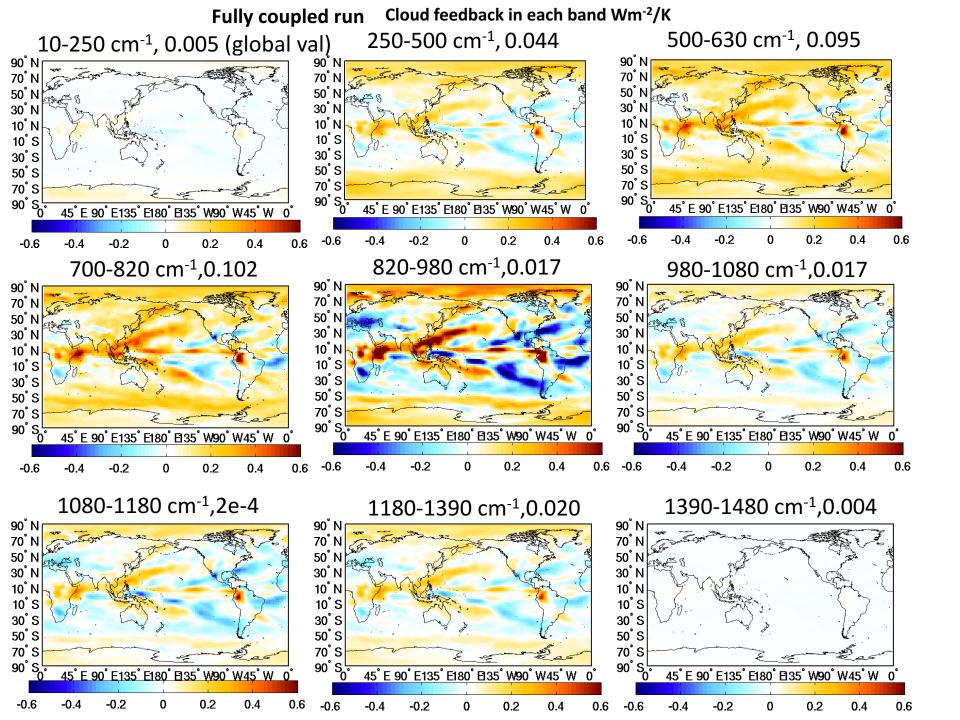
Differences of Model - Obs



Band-by-band LW cloud feedback in the NCAR CESM (2xCO₂ run)

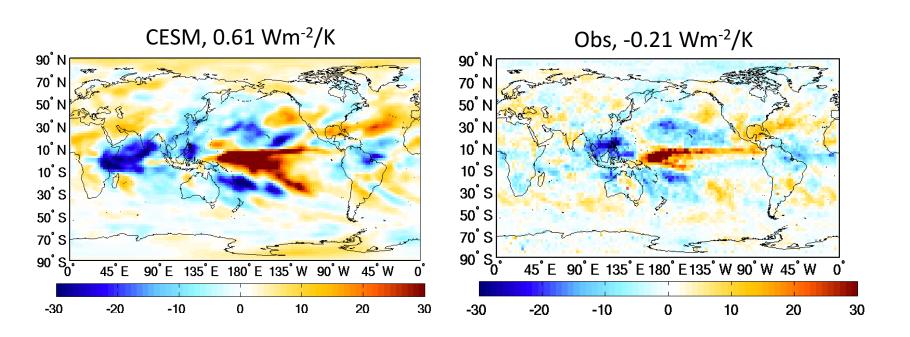


Broadband LW cloud feedback Slab ocean run: 0.25 Wm⁻²/K Fully coupled run: 0.31 Wm⁻²/K

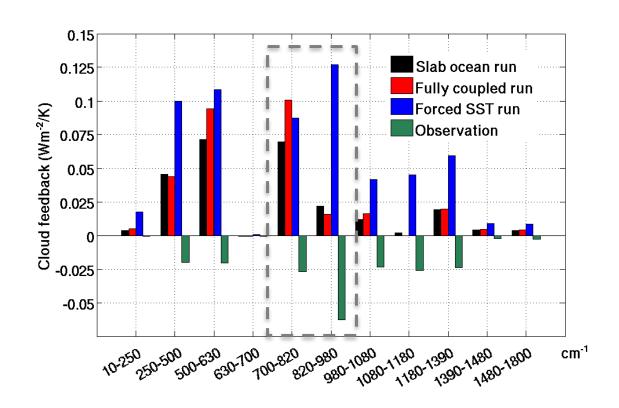


Short-term fluctuation of 2003-2015 (Preliminary)

- CESM simulation: using Dessler's method to obtain an estimation of short-term cloud feedback
- Observation: applying Yue et al. (2016) to MODIS,
 AIRS and CERES data to obtain the same quantity



Long-term vs. short-term contrast



Broadband LW cloud feedback Slab ocean run: 0.25 Wm⁻²/K

Fully coupled run: 0.31 Wm⁻²/K

Forced SST run: 0.61 Wm⁻²/K

Observation: -0.21 Wm⁻²/K

Conclusion and Discussion

- Spectral decomposition helps revealing compensating biases.
 - For cloud, it helps untangle the biases in CTH and cloud amount
- For RRTMG bandwidths, 820-980cm⁻¹ contributes most to the CRE. But 700-820 cm⁻¹ contributes most to the cloud feedback (2xCO₂ run)
- The long-term vs. short-term cloud feedback has different spectral decomposition
 - Implications

Spectral Flux Geophysical variables Broadband Radiation Budget $F_{v} = \int_{0}^{2\pi} d\phi \int_{0}^{\frac{\pi}{2}} I_{TOA}(v; \theta, \phi) \cos\theta \sin\theta d\theta$ $F = \int_{\Lambda v} F_{\underline{v}} dv$ T(z)Spectral Radiances $q_{H2O}(z)$ $I_{TOA}(v;\theta,\phi)$ $q_{O3}(z)$ **Spectral** $q_{CH4}(z)$ Radiative **Broadband Radiative Feedbacks Feedbacks** $\lambda_{x} = -\frac{\delta_{x} F}{\delta X} \frac{\delta X}{\delta T_{s}}$ $\lambda_{x_{v}} = -\frac{\delta_{x} \overline{F_{v}}}{\delta X} \frac{\delta X}{\delta T_{c}}$ **Aerosols** T_{skin} , $\epsilon_s(v)$ Cloud, **ISCCP** effort Energy budget and feedbacks Sounding community community

Thank You!

References:

- Huang et al., 2008: Spectrally resolved fluxes derived from collocated AIRS and CERES measurements and their application in model evaluation, Part I: clear sky over the tropical oceans, JGR-Atmospheres, 113, D09110, doi:10.1029/2007JD009219.
- 2. Chen et al., 2013: Comparisons of clear-sky outgoing far-IR flux inferred from satellite observations and computed from three most recent reanalysis products, *Journal of Climate*, 26(2), 478-494, doi:10.1175/JCLI-D-12-00212.1.
- 3. Huang et al., 2014: A global climatology of outgoing longwave spectral cloud radiative effect and associated effective cloud properties, *Journal of Climate*, 27, 7475-7492, doi:10.1175/JCLI-D-13-00663.1.
- 4. Huang, X. L., X. H. Chen, B. J. Soden, X. Liu, 2014: The spectral dimension of longwave feedbacks in the CMIP3 and CMIP5 experiments, *Geophysical Research Letters*, 41, doi:10.1002/2014GL061938.
- 5. Yue, Q., B. H. Kahn, E. J. Fetzer, M. Schreier, S. Wong, X. H. Chen, X. L. Huang, 2016: Observation-based Longwave Cloud Radiative Kernels Derived from the A-Train, Journal of Climate, 29, 2023-2040.

Monthly gridded spectral flux and CRE available via http://www-personal.umich.edu/~xianglei/datasets.html.

The spectral radiative kernels available upon request.

Backup slides

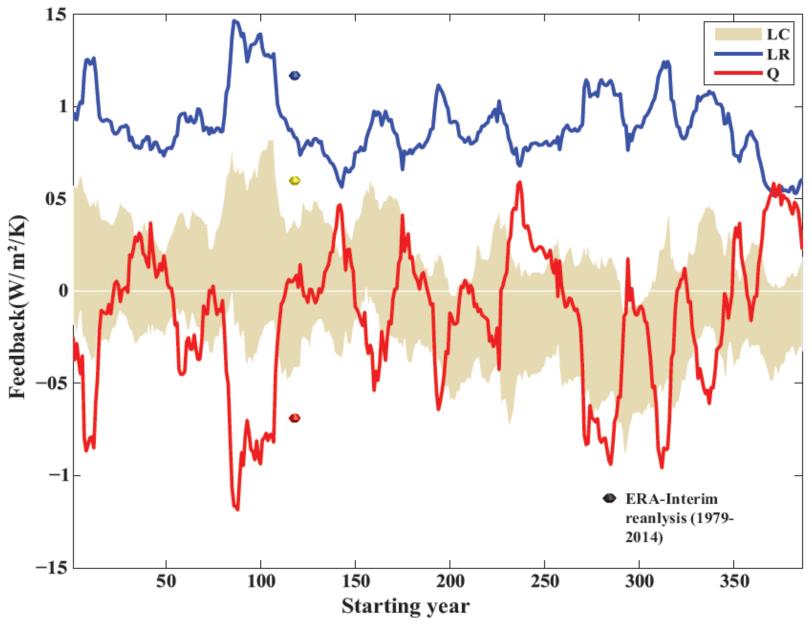


Fig 3. Short-term longwave cloud feedback (LC), lapse-rate feedback (LR), and water vapor feedback (Q) derived from different segments of 35-year CM3 simulations.



All collocated clear-sky observations in 2004 (80 $^{\circ}$ S-80 $^{\circ}$ N)

	Daytime	Nighttime	
Surface Type	OLR _{AIRS_Huang} -OLR _{CERES} (Wm ⁻²)	OLR _{AIRS_Huang} - OLR _{CERES} (Wm ⁻²)	
Forest	0.58 ± 1.43	-0.42 ± 1.41	
Savannas	-0.03 ± 2.52	0.68 ± 1.50	
Grasslands	0.19 ± 2.61	0.63 ± 1.65	
Dark Desert	-0.71 ± 2.85	0.36 ± 1.74	
Bright Desert	1.67 ± 2.62	1.42 ± 2.28	
Ocean	1.09 ± 1.55	0.90 ± 1.26	

(Chen et al., J Climate, 2013)

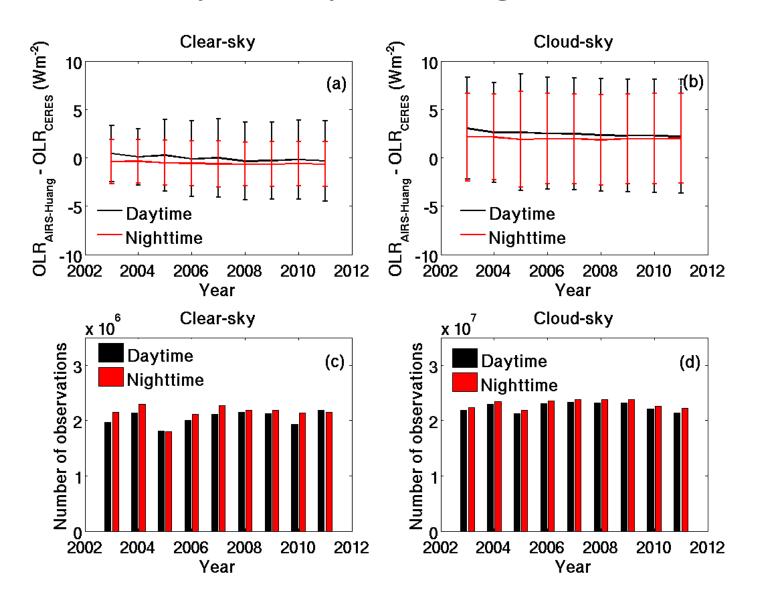
CERES 2σ radiometric calibration uncertainty: 1% (i.e. ~ 2.5W m⁻²)

Stratifying OLR_{AIRS_Huang}-OLR_{CERES} (Wm⁻²): cloudy observations over the lands

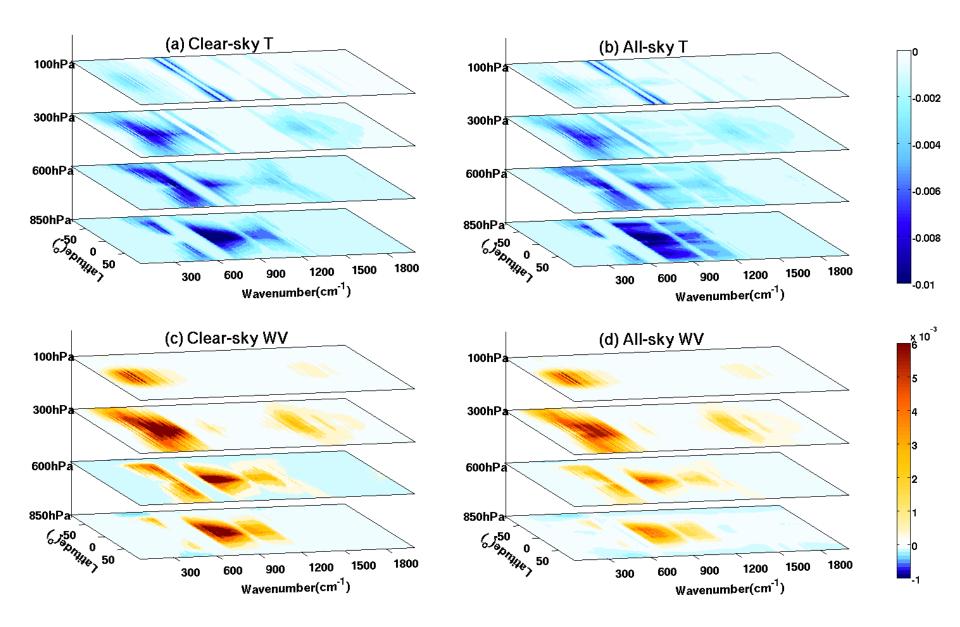
	Over deserts			Over non-desert lands		
f	<15k	15K-40K	>40K	<15k	15K-40K	>40K
0.001-0.5	2.44±3.79	3.25±5.12	1.49±7.61	2.34±2.86	3.62±4.48	2.84±5.94
	(0.9%)	(1.2%)	(0.5%)	(0.8%)	(1.3%)	(1.0%)
0.5.0.75	2.79±4.16	3.34±7.80	1.39±12.75	2.90±3.86	4.24±7.25	2.61±11.38
0.5-0.75	(1.1%)	(1.3%)	(0.5%)	(1.1%)	(1.7%)	(1.0%)
0.75-0.999	2.67±3.67	1.45±6.47	-1.17±10.97	2.81±3.56	3.14±6.68	0.47±11.45
	(1.1%)	(0.6%)	(-0.5%)	(1.2%)	(1.4%)	(0.2%)
0.999-1.0	2.61±2.80	3.15±4.00	1.28±6.64	2.86±2.83	4.04±4.33	2.48±7.16
	(1.2%)	(1.6%)	(0.7%)	(1.3%)	(2.0%)	(1.5%)

CERES 2σ radiometric calibration uncertainty: 1% (i.e. ~ 2.5W m⁻²)

Global OLR_{AIRS_Huang}-OLR_{CERES}: annual means and year to year changes

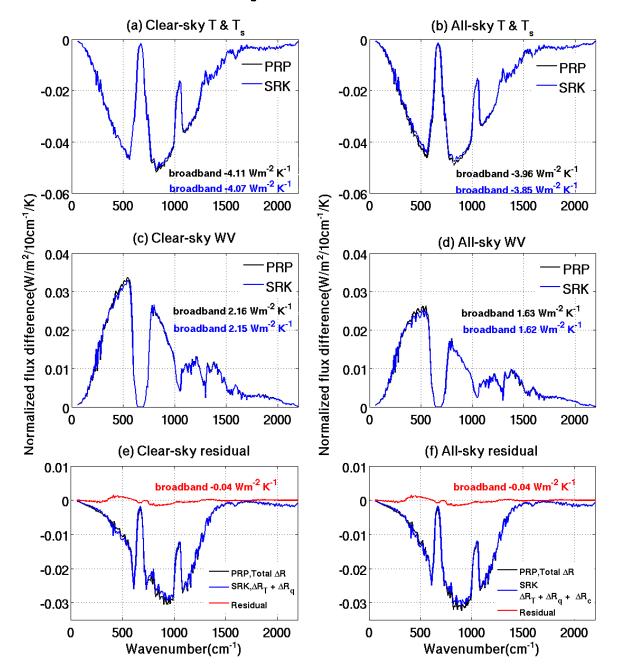


Construction of the SRK





Validation: comparisons with the PRP results





Validation: comparisons with the PRP results

